Application of Geophysical Techniques in Artificial Recharge to Ground Water



K.C.MONDAL Scientist-B Mobile no-09753079567

E-mail: mondalkamalesh13@gmail.com

Investigations/Studies Required

- Topographical surveys
- Hydrometeorological Studies
 - In-situ precipitation, rainfall pattern, evaporation losses and climatologically features etc.
- Hydrological Studies
 - Runoff pattern, drainage pattern etc.
- Soil Infiltration Studies
- Geochemical Studies

?

- Hydrogeological Studies
- Geophysical Investigations
- Engineering Design and Construction

Please remember this 8 words

If fail to Plan, you plan to fail"

Types of rocks

Three Types

- 1. Igneous rock
- 2. Sedimentary rock
- 3. Metamorphic rock

Total water on earth

→ 3% (37.5 mill. Cu.Kms) Fresh Water

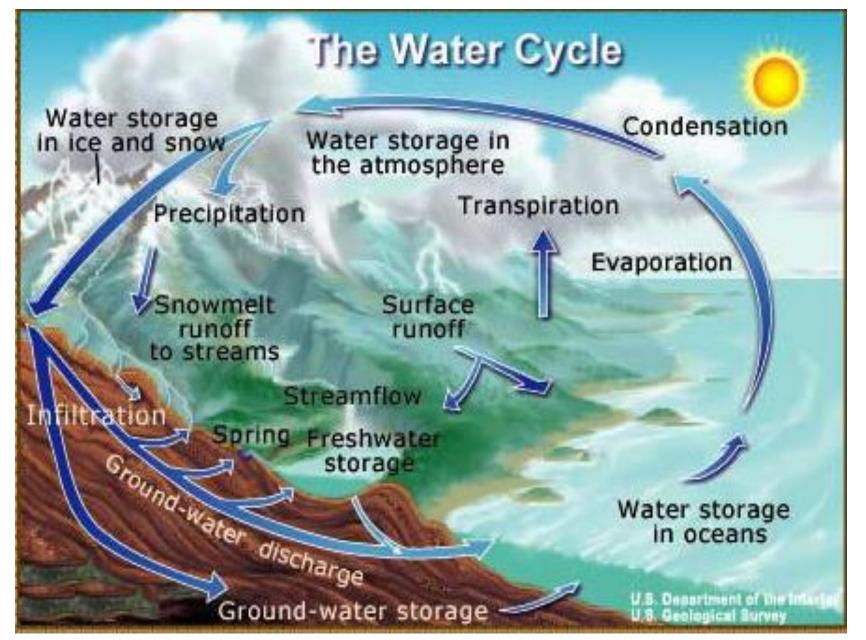
→ 97% (1320 mill. Cu. Kms.) (Saline Water)

Distribution of fresh water on earth

11% (ground water up to 800m) 11% (ground water below 800m) 277%(in glacier & ice caps) 800m)

11% of the total fresh water on earth (up to 800 m depth) is available for extraction

Water: Where does it come from ?



Hydrogeology

(*hydro-* meaning water, and *-geology* meaning the study of the Earth) is the area of geology that deals with the distribution and movement of groundwater in the soil and rocks of the Earth's crusts, (commonly in aquifers).

What is Ground water ?

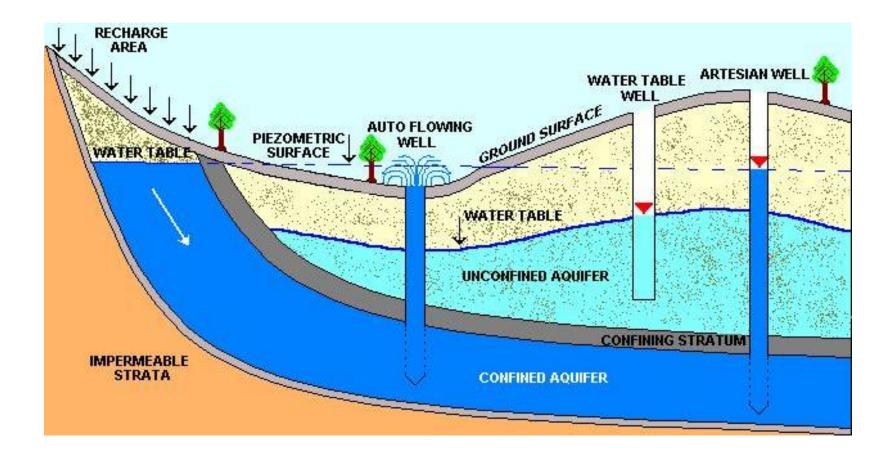
All subsurface water that fills the pores, fractures, and other spaces in rock strata below the Water Table(in the saturated zone of geologic formations).

- Aquifer: A saturated permeable geological unit that is permeable enough to yield economic quantities of water to wells. (Sand)
- Aquitard: A geological unit that is permeable enough to transmit water in significant quantities when viewed over large areas and long periods, but its permeability is not sufficient to justify production wells being placed in it.(Arenacious Shale & Clay)
- Aquiclude: A geological unit which can only store water but can't transmit significant amount.(Shale & Clay)
- Aquifuge: A geological unit with no interconnected pores and hence can neither absorb nor transmit water(Massive Granite/Marble/Basalt)

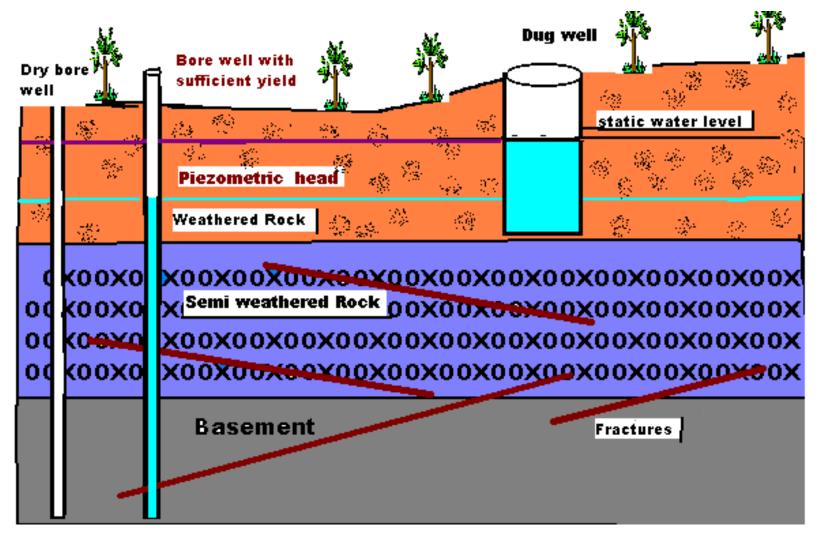
In Surface Geophysical Survey we can consider the formation are of two types depending their compactness

- 1. Hard Formation &
- 2. Soft formation

Aquifers in Alluvium Formation



Occurrence of Ground Water in Hard Formation



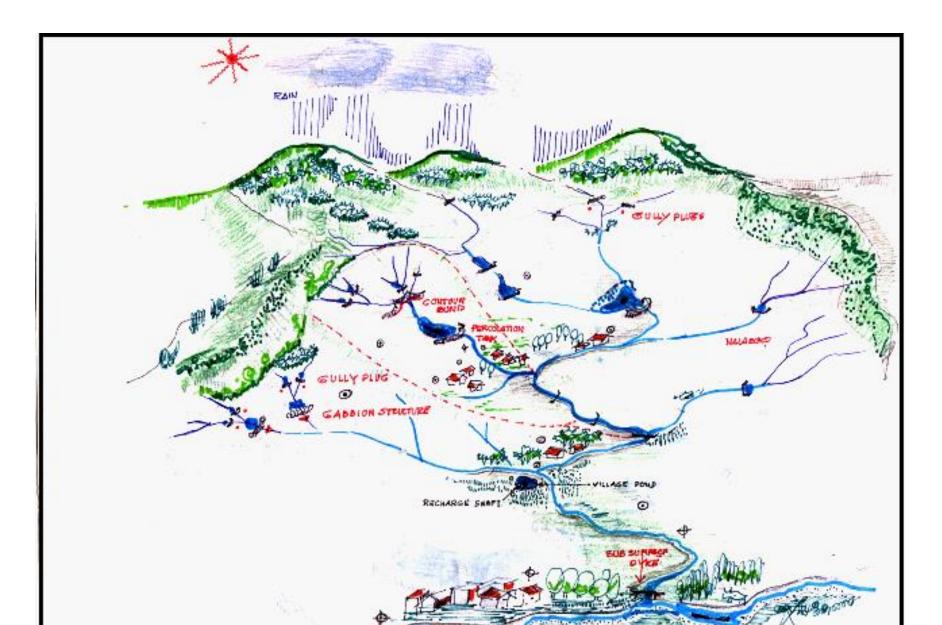
What is Geophysics ?

Measurements of earth's physical parameters at the earth surface.

Objectives of Geophysical Surveys

- The prime objective of geophysical survey is to unravel the hidden sub-surface hydrogeological conditions to save money and time.
- The other major objectives include
 - Delineation of sub-surface water bearing horizons.
 - Identification of fractured zones in hard rock areas.
 - Estimation of overburden thickness and delineation of bed rock topography.
 - Mapping of aquifer geometry.
 - Delineation of fresh/saline groundwater interface.
 - Identification of suitable areas/locations for artificial recharge of groundwater.

Artificial Recharge Structure



Information's Required to select site, zone or area for artificial recharge

- Lithologic variations, thickness, depth and lateral extent of the vadose zone and the aquifers
- Water level / Peizometric head
- Obstacles and direction of groundwater movement
- Groundwater quality and hydrogeologic parameters

Geophysical Objectives in Different Terrain for AR

Hard Rocks: in the weathered zone and fracture zones or maintain water level in it through constraining the movement of groundwater

Limestones : in solution channels or cavities after identifying their location and path of recharge

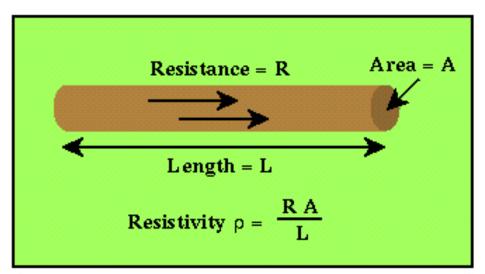
Alluvium: in near surface sandy soil, sand zone, flood plain, stream bed, sand dunes, . Also in confined aquifers through injection wells

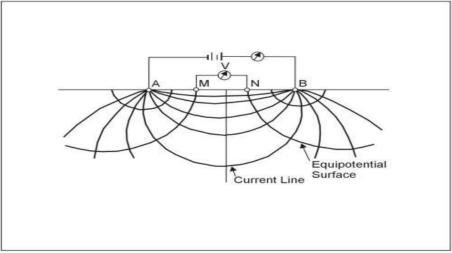
Urban areas: in open areas/porous pavement/parking place

Electrical Resistivity Surveys

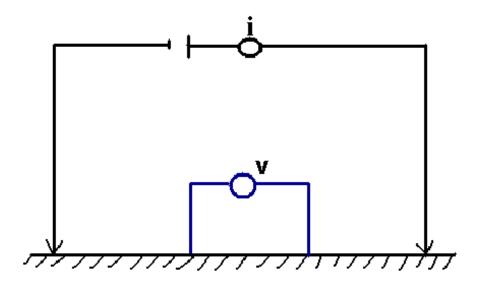
- Started by American Scientist *SCHLUMBERGER* during year 1912 -1914
- It is based on Ohm's Law (V=IR)
- Resistivity of a material is defined as the resistance un ohm between opposite faces of a unit cube (l=1 m, A= $1m^2$) of the material
- If the measurement is made over a semi infinite space of homogeneous and isotropic material, then it will be a true resistivity
- If medium is inhomogeneous or anisotropic then it will be a apparent resistivity

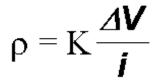
Electrical Resistivity Methods Principle – Ohm's Law





In simplest form of resistivity method, a known amount of electric current is sent into the ground through a pair of electrodes and potentials developed due to current within the ground are measured across another pair of electrodes on the ground.





Advantages of Electrical Resistivity Survey

Relatively direct and simple response to the groundwater conditions

Low cost of survey

Easy field data interpretation procedures

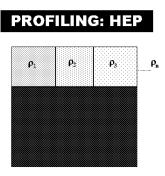
APPLICATION OF GEOPHYSICAL TECHNIQUES IN GROUND WATER EXPLORATION

- Narrow down the target area.
- Pinpoint the probable site for ground water exploration.
- Comparative picture of the sub-surface litho environment.
- Identification of brackish/fresh ground water interface.
- Direction of ground water flow under natural/artificial recharge processes.

Types of Electrical Resistivity Survey

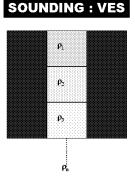
Horizontal Electrical Profiling (HEP)

It measures lateral variations in earth resistivity

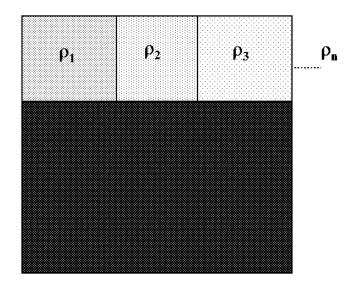


Vertical Electrical Sounding (VES)

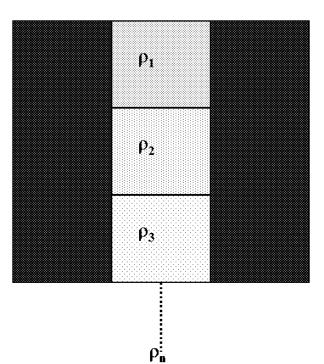
It measures resistivity variations with depth



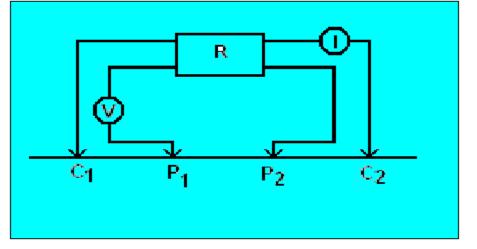
PROFILING: HEP



SOUNDING : VES



1. Wenner Configuration:



The formula for measuring the electrical Resistivity (apparent) is: $\rho_a = 2 \pi a R,$ where a = Distance between two successive electrodes R = V/I (V= Voltage, I = Current)

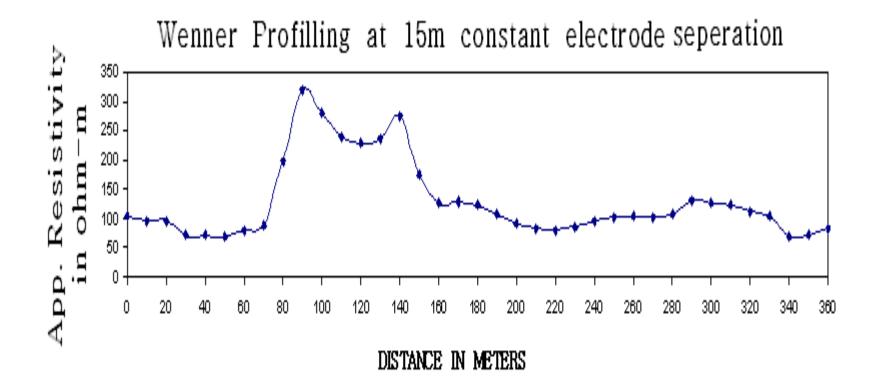
The unit for measuring apparent Resistivity is ohm-metres. In this method all 4 electrodes are moved constantly after each reading and the depth of penetration is about 1/3 of the total electrode separation, (C1-C2).

The Wenner array itself has got three variations, depending the change of positions of electrodes. These are,

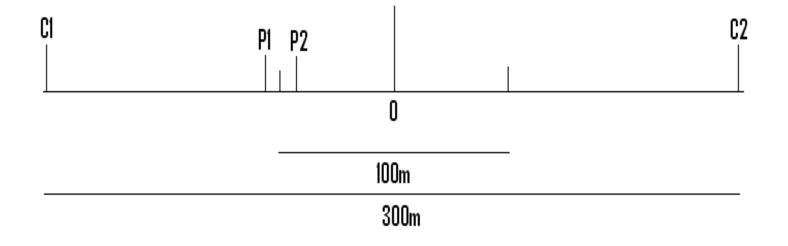
 α - Wenner, where the electrode positions are C1, P1, P2, C2. This is also called normal Wenner array.

- β Wenner, where the electrode positions are C1, C2, P1, P2.
- γ Wenner, where the electrode positions are C1, P1, C2, P2.

		:	Surface Geop	hysical Su	rvey		
				Profilling			
	Rajiv	Gandhi Na	ational Ground		ing Instit	tute, Raipur	
			ent of India, Mi				
			Development				
Village : Geophysicist :			Village : Geophysicist :				
Block :			Block :				
District :				District : Direction:			
Date:			_	Date:			
Station	К	R	App. Rest.	Station	К	R	App. Rest.
1				1			
2				2			
3				3			
4				4			
5				5			
6				6			
7				7			
8				8			
9				9			
10				10			
11				11			
12				12			
13				13			
14				14			
15				15			
16				16			
17				17			
18				18			
19				19			
20			1	20			
21				21			
22				22			
23				23			
24				24			
25				25			
26				26			
27				27			
28				28			
29			1	29			
30				30			



Layout of Gradient Resistivity Profiling(GRP)



Gradient Resistivity Profiling(GRP

Formula

$$\rho_{a} = \frac{\pi}{2I} \times \frac{L^{2}(1-x^{2}/L^{2})^{2}}{1+x^{2}/L^{2}} \times \frac{\Delta V}{I} = K \frac{\Delta V}{I} = K \times R$$

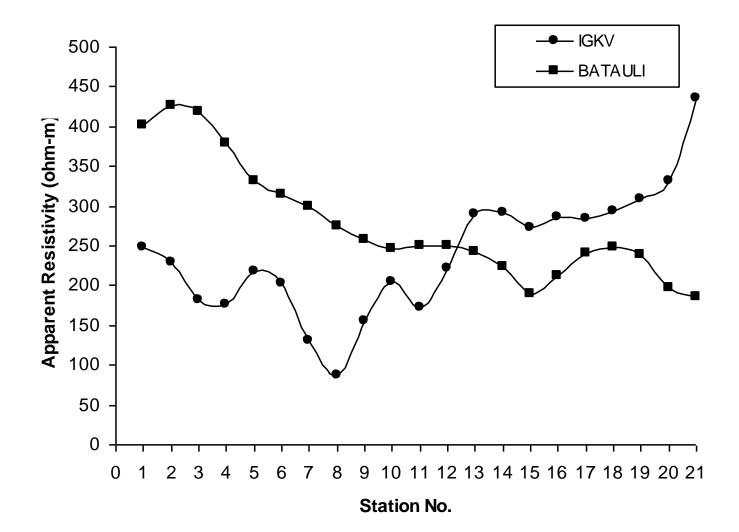
P_a = Apparent Resistivity

- **2L** = Distance between two current electrode
- **2I** = Distance between potential electrodes

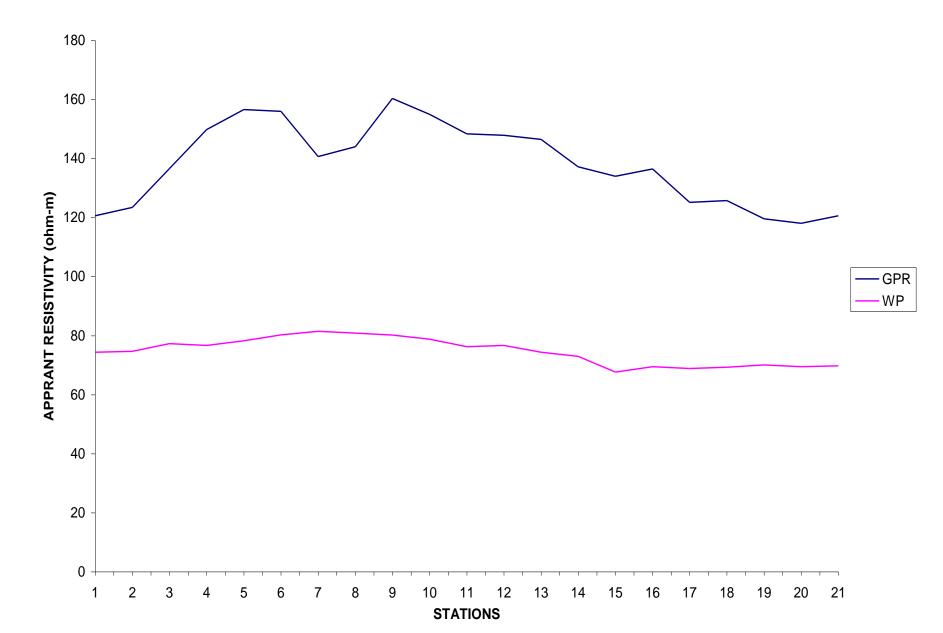
X = Distance between the Centre of the current electrodes to the centre of the potential electrodes

R = Resistance of the formation

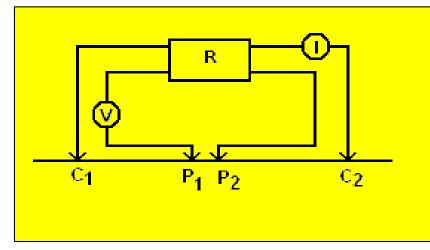
	Surface Geophysical Survey									
	Gradient Resistivity Profiling(GRP)									
Rajiv Gandhi National Ground Water Training Institute, Raipur										
Government of India, Ministry of Water Resources, River Development & Ganga Rejuvenation										
Village :	Village : Geophysicist :		Village :		Geophysicist :					
Block :		Profiling No:		Block :		Profiling No:				
District :	District : Direction:			District :		Direction				
Date:	Date:					Date:				
K for AB=30	00, MN=10n	n& Stn. interval=5m		K for AB=300, MN=10m& Stn. interval=5m						
Station	к	R	App. Rest.	Station	к	R	App. Rest.			
1	5026.8	8		1	5026.88					
2	5370.5	3		2	5370.53					
3	5694.4	8		3	5694.48					
4	5993.9	3		4	5993.93					
5	6264.2	7		5	6264.27					
6	6501.1	9		6	6501.19					
7	6700.8	1		7	6700.81					
8	6859.7	8		8	6859.78					
9	6975.3	5		9	6975.35					
10	7045.5	2		10	7045.52					
11	7069.0	5		11	7069.05					
12	7045.5	2		12	7045.52					
13	6975.3	5		13	6975.35					
14	6859.7	8		14	6859.78					
15	6700.8	1		15	6700.81					
16	6501.1	9		16	6501.19					
17	6264.2	7		17	6264.27					
18	5993.9	3		18	5993.93					
19	5694.4	8		19	5694.48					
20	5370.5	3		20	5370.53					
21	5026.8	8		21	5026.88					



Gradient Resistivity Profiling At IGKV (Raipur) and Batauli (Surguja)



2. Schlumberger Configuration:



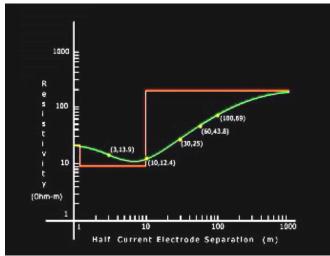
The formula for measuring the electrical resistivity (apparent) is

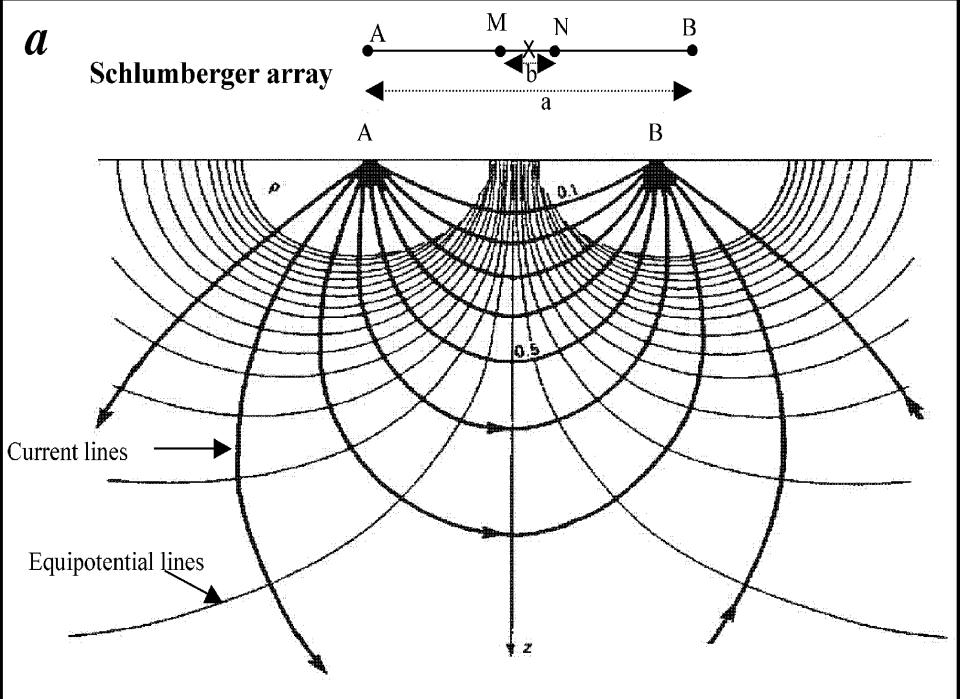
 $\rho_{a} = \pi R - \frac{[(AB/2)^{2} - (MN/2)^{2}]}{MN}$

Where AB/2 = Half current electrode separation

MN/2 = Half potential electrode separation

R = V/I (V= Voltage) (I= Current)

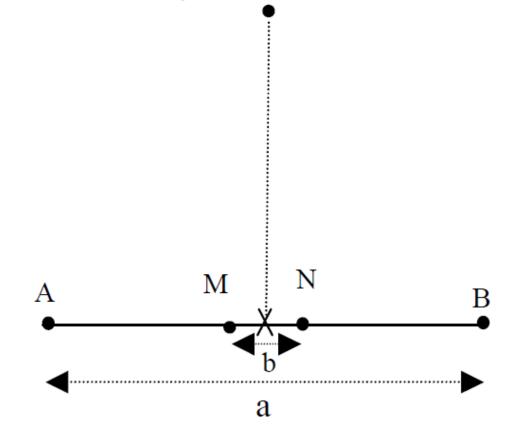


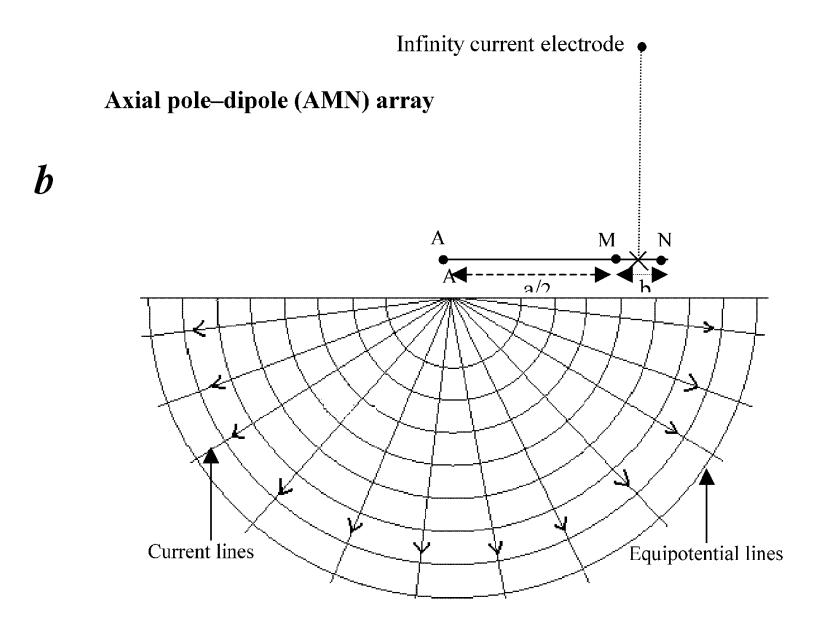


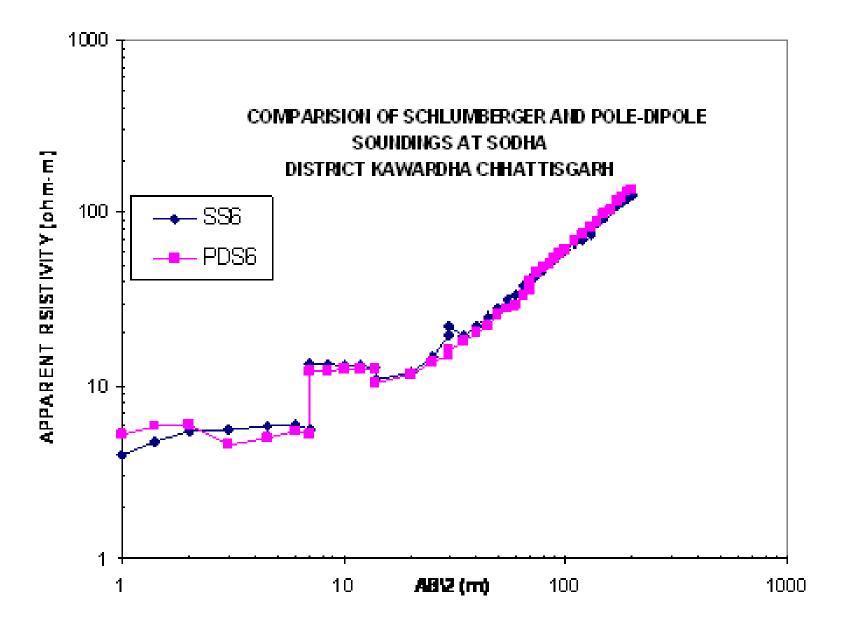
Surface Geophysical Survey								
	Vertical Electrical Sounding (VES)							
Rajiv Ga	Rajiv Gandhi National Ground Water Training & Research Institute, Raipur							
Government of India, Ministry of Water Resources,								
River Development & Ganga Rejuvenation								
Village :			Geophysicist :		Date :			
Block :				VES No. :				
District :			Survey No. :	VES Direction :				
S.No.	AB/2	MN/2	Geometrical	Resistance	Apparent			
	(m)	(m)	Factor (K)	(ohm)	Resistivity			
					(ohm-m)			
1	2	0.4	15.09					
2	3	0.4	34.73					
3	4.5	0.4	78.92					
4	6	0.4	140.80					
5	8	0.4	250.80					
6	10	0.4	392.22					
7	10	2	75.43					
8	12	0.4	565.08					
9	15	2	173.64					
10	20	2	311.14					
11	25	2	487.92					
12	30	2	703.99					
13	30	5	275.00					
14	35	5	377.14					
15	40	5	494.99					
16	45	5	628.56					
17	50	5	777.84					
18	60	5	1123.55					
19	70	5	1532.12					
20	70	10	754.27					
21	80	10	989.98					
22	90	10	1257.12					
23	100	10	1555.69					
24	120	10	2247.10					
26	150	10	3519.94					
27	200	10	6269.89					
28	250	10	9805.54					
29	300	10	14126.89					
30	300	20	7039.87					
31	350	20	9593.40					
32	400	20	12539.77					
33 500 20 19611.07								

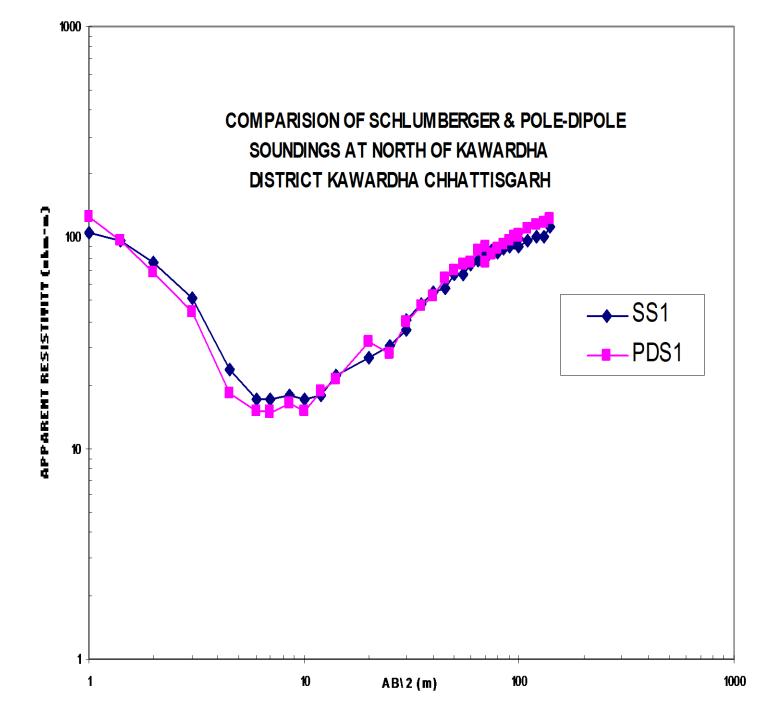
Pole-Di-Pole (Half Schlumberger Method)

Infinity current electrode









Interpretation of VES Data

- Plot of apparent Resistivity values for different electrode separation on Logarithmic scale
- Interpretation through curve matching techniques
- Interpretation through Computer Software
 - Forward modeling
 - Inverse modeling

Interpretation can be done by two methods

- **1. Manual Data Interpretation**
- 2. Computerised Data Interpretation

Interpretation of VES Data

Interpretation can be done by two methods

1. Manual Data Interpretation

2. Computerised Data Interpretation

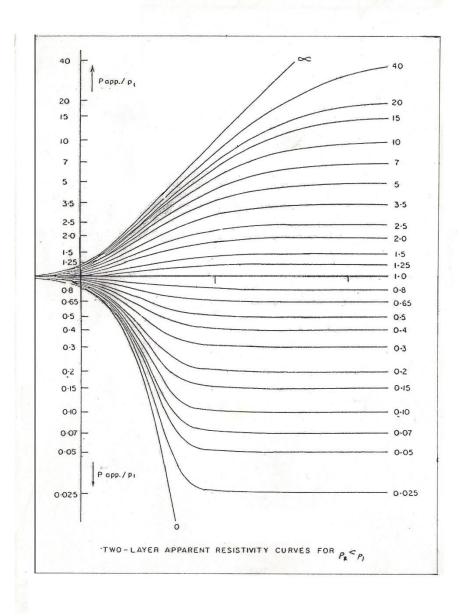
Type of VES curves

- $H type \ (\rho_1 > \rho_2 < \rho_3)$
- K type ($\rho_1 < \rho_2 > \rho_3$)
- A type ($\rho_1 < \rho_2 < \rho_3$)
- Q type ($\rho_1 > \rho_2 > \rho_3$)



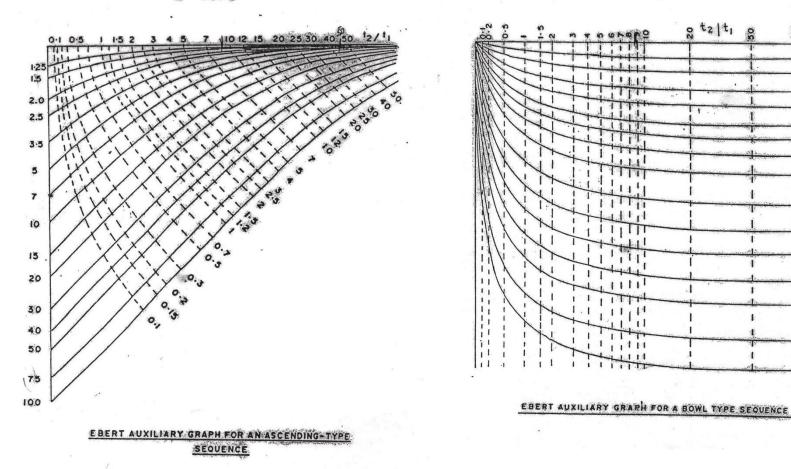
Two Layer Master Curve

Multilayer problem is reduced to the simple two layer case with four charts of auxiliary curves.



Auxiliary Curve

A - TYPE



"H'- TYPE

St2|t1

.1

1

11

.

1

1

100

0.8

0.6

0.5

0.4

0.3 0.25

0.20

0.15 P2

- 0.07

0.05

0.035

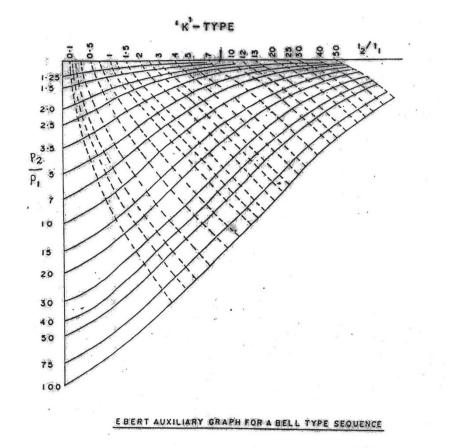
- 0.025

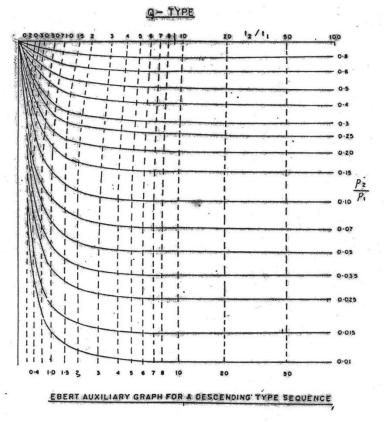
0.015

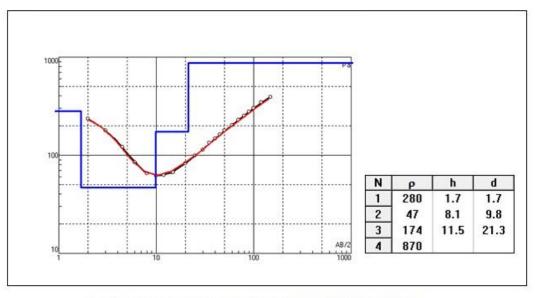
0.01

0.10 P1

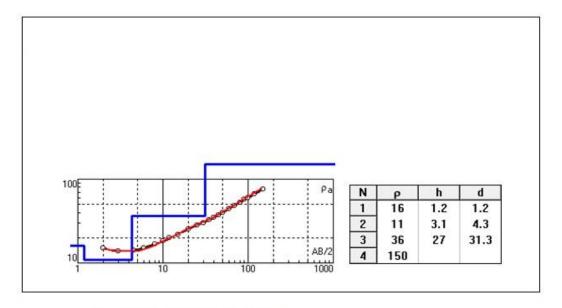
Auxiliary Curve



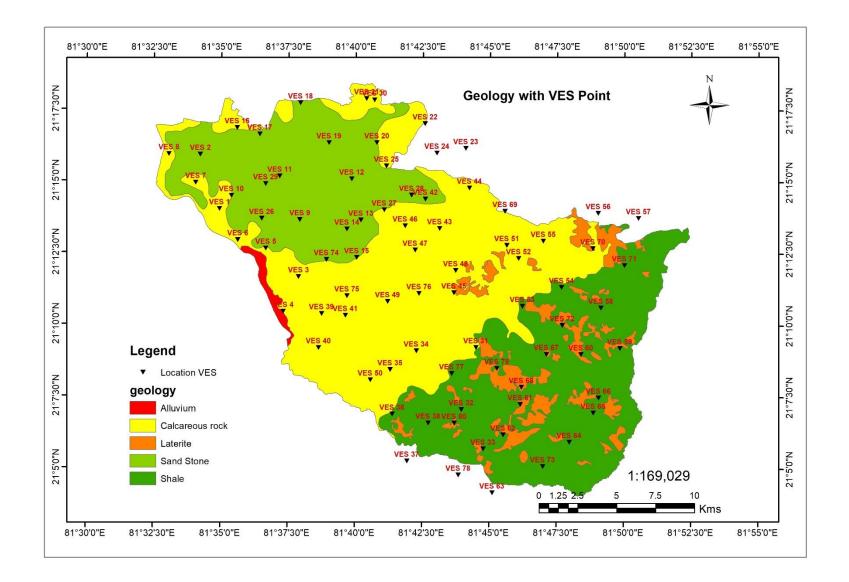


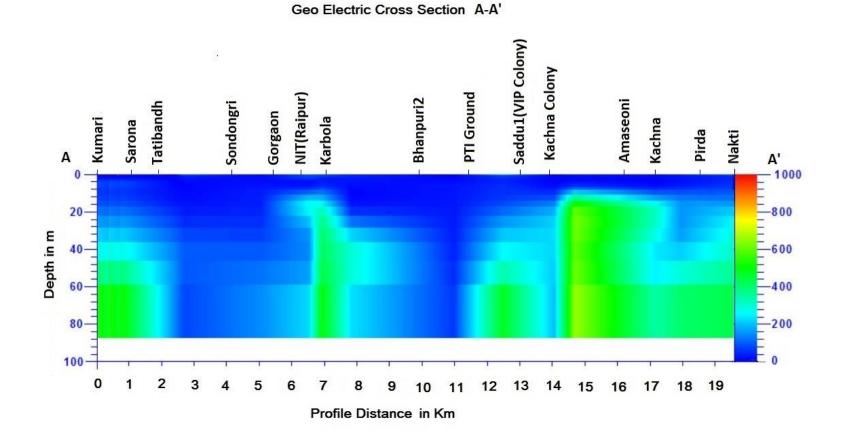


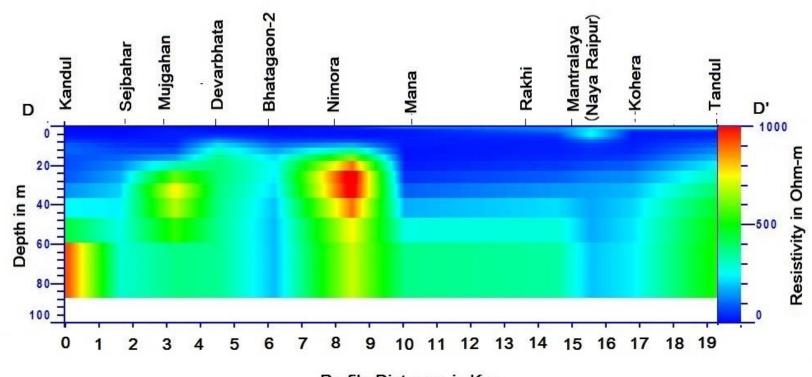
Field Curves and Interpreted Results of VES-15 (Kamal Vihar)



Field Curves and Interpreted Results of VES-16 (Sondongri)

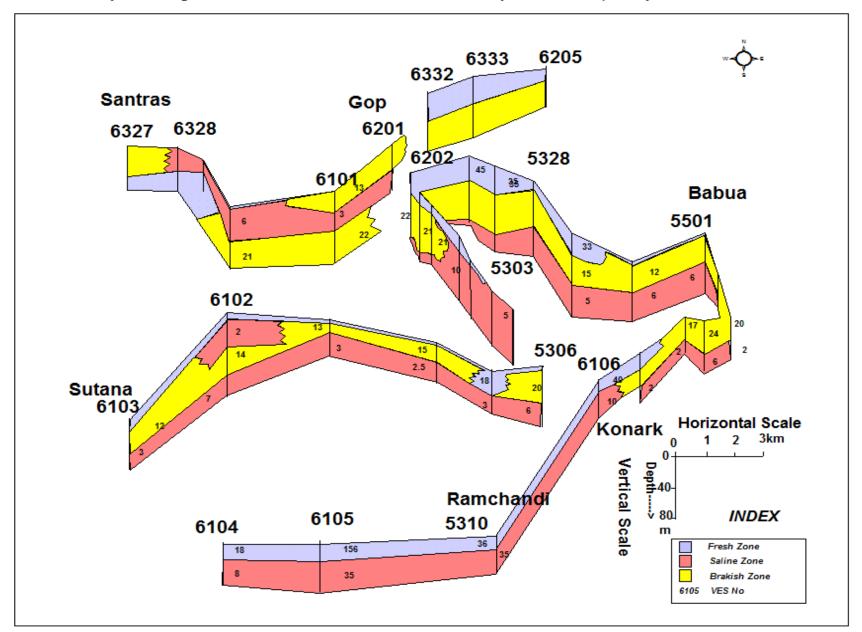




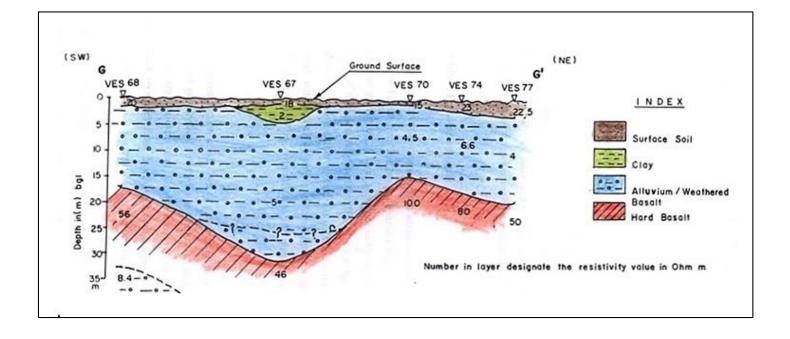


Geo Electric Cross Section D-D'

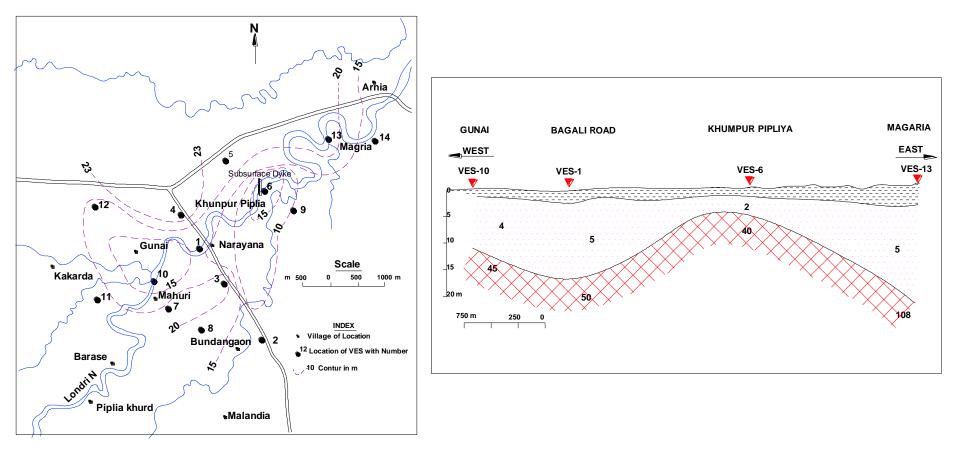
Profile Distance in Km



Geo-electrical Cross Section along Percolation Dam of Sikheri , District- Mandasoure (MP)



Londhri River bed, Dewas district (MP)



Any doubt?

